

Use of Stress Drop Models to Interpret Geodetic Measurements at Loma Prieta

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Objective

Surface deformation due to slip in the earth's crust is typically interpreted using the well-known dislocation method. Here, we employ an alternative approach in which slip regions are approximated by planar zones of prescribed stress drop (rather than relative displacement, i.e., dislocation). By applying this approach to the observed coseismic geodetic data associated with the 1989 Loma Prieta, California, earthquake, we are able to obtain estimates for the slip induced stress drop and moment, and the critical energy release rate at the termination of rupture.

Results

The Loma Prieta event is modeled as shear slip of an inclined elliptical crack embedded in an elastic half-space. The shear stress drop is assumed to be uniform with a component $\Delta\tau_s$ along the strike and a component $\Delta\tau_d$ along the dip direction of the fault. The resulting slip is calculated using the method of Lee et al. [1987]. Normal relative displacement over the fault plane, which tends to occur because of the free surface, is constrained to be zero. A fairly large set of forward searches was performed to find the model parameters that give a best fit to the observed coseismic geodetic data [Lisowski, personal communication, 1990] subject to the constraint that the fault geometry agrees roughly with the aftershock locations [Dietz and Ellsworth, 1990].

Reasonable agreement with the deformation data is obtained with a model for which subequal components of stress drop along strike and dip are applied over an oblique fault plane. Table 1 shows the fit of the best model to the geodetic data. For brevity, the data with magnitudes smaller than 35 mm are not shown. The mean misfit to all data is 40.8 mm, about 2.5 times the data error which is about 15 mm. The largest misfit occurs for the line lp1-lp2. As pointed out by Lisowski et al. [1990], it is likely that this line is affected by local movements, given the widespread surface cracking and secondary faulting observed in the epicentral area [U.S.G.S. Staff, 1990]. The strike, the inclination angle (downdip), the depth of the fault center, the fault length, and the fault width are about N48°W, 75°, 11.5 km, 50 km, and 17 km, respectively. The stress drop components determined by minimizing the difference between the observed and predicted surface deformation are $\Delta\tau_s = 1.2$ MPa and $\Delta\tau_d = 1.5$

MPa. The calculated geodetic moment is 2.5×10^{19} Nm. The maximum energy release rate (energy released per unit area of fault advance) occurs near the top of the fault and is estimated to be 5.5×10^6 J/m². If it is assumed that the slip propagated according to the criterion that the energy release rate is equal to a critical value, then the estimated value is the critical energy release rate corresponding to arrest of the earthquake. The slip distribution over the fault surface is approximately elliptical. The average slips in the strike and (reverse) dip directions are $[\bar{u}_s] = 1.1$ m and $[\bar{u}_d] = 0.9$ m, respectively.

The geodetic moment, the dipping angle, and the ratio between strike and dip slip are very close to those (2.8×10^{19} Nm, 70° , and 1.3) of Lisowski et al. [1990]. The magnitudes of our strike and dip slip are smaller than theirs (1.6 m and 1.2 m) because our best model prefers a fault geometry having a larger area. If we use the geometry suggested by Lisowski et al., the mean misfit to data is increased to be 64 mm.

References

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Publication

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Table 1. Model Fit to Geodetic Data

Station - Station	Observed (mm)	Calculated (mm)	Residual (mm)
<u>Geodolite Line Length Changes</u>			
allison bmt rf	-39.2 ± 11.8	-25.1	-14.1
american hamilton	75.0 ± 9.4	52.7	22.3
american loma use	-66.5 ± 7.8	-32.8	-33.7
biel eagle rk	101.5 ± 7.2	60.3	41.2
biel loma use	235.4 ± 11.6	218.4	17.0
biel mindego	-107.5 ± 8.8	-98.1	-9.4
bmt rf loma use	176.2 ± 18.1	106.9	69.3
butano dump	-39.7 ± 5.7	-32.3	-7.4
butano eagle rk	-95.2 ± 5.5	-56.2	-39.0
butano pom	-46.2 ± 11.1	-32.9	-13.3
gilroy llagas	-48.5 ± 7.9	-50.8	2.3
hamil ec llagas	37.3 ± 8.9	3.4	33.9
hamil ec sheeprm2	-42.1 ± 12.7	-24.9	-17.2
llagas lp1	-54.0 ± 6.0	5.0	-59.0
llagas sheeprm2	74.3 ± 11.9	14.6	59.7
loma use mindego	175.0 ± 12.8	122.9	52.1
allison lomancer	109.0 ± 8.0	78.1	30.9
eagle rk lomancer	259.7 ± 5.8	202.5	57.2
hamilton lomancer	51.1 ± 6.0	52.6	-1.5
lomadwr pr6	-274.1 ± 20.1	-236.9	-37.2
lp1 lp2	-40.9 ± 5.2	146.4	-187.3
lp1 lp4	-212.0 ± 7.9	-221.1	9.1
brush 2 fremont	-18.1 ± 10.0	-9.7	-8.4
brush 2 mulligan	49.4 ± 6.9	30.1	19.3
chamber vargo	123.9 ± 7.2	127.9	-4.0
fremont juan	42.5 ± 6.0	6.3	36.2
gilroy juan	-63.4 ± 11.4	-33.8	-29.6
juan mulligan	-36.8 ± 8.7	-13.2	-23.6
lp1 vargo	-204.0 ± 15.3	-239.0	35.0
fremont sargent	54.2 ± 12.9	-15.6	69.8
canada fairview	-39.3 ± 12.2	-8.0	-31.3
canada sargent	-86.3 ± 25.0	-1.4	-84.9
fairview gilroy	-68.2 ± 13.7	-19.3	-48.9
gilroy sargent	-89.9 ± 12.6	-16.1	-73.8
gilroy sheeprm2	-69.8 ± 11.1	-19.0	-50.8
canada gilroy	-51.9 ± 7.5	-8.7	-43.2
p2 p3	44.6 ± 23.1	-5.5	50.1
p3 p4	66.6 ± 14.0	5.2	61.4
p3 p5	66.6 ± 23.1	11.3	55.3
p4 p5	46.0 ± 11.3	6.3	39.7

Table 1. (continued)

p5	p6	87.3 ± 13.7	-0.8	88.1
p5	p7	-43.3 ± 13.5	-9.6	-33.7
p7	p8	73.7 ± 10.6	8.0	65.7
p5	chamb 2	-41.8 ± 12.0	-7.3	-34.5
<u>GPS East</u>				
lp1	allison	34.8 ± 5.9	70.2	-35.4
lp1	eagle un	225.2 ± 11.8	243.6	-18.4
lp1	hamilton	65.6 ± 7.0	66.0	-0.4
lp1	brush 2	41.3 ± 15.6	78.4	-37.1
<u>GPS North</u>				
lp1	allison	-110.6 ± 3.8	-103.7	-6.9
lp1	eagle un	-205.4 ± 5.6	-145.3	-60.1
lp1	hamilton	-106.6 ± 2.6	-96.2	-10.4
lp1	brush 2	-229.2 ± 4.6	-205.1	-24.1
<u>GPS Up</u>				
lp1	allison	-105.5 ± 33.4	-139.1	33.6
lp1	eagle un	-211.8 ± 24.7	-202.6	-9.2
lp1	hamilton	-114.4 ± 25.5	-131.4	17.0
lp1	brush 2	-40.0 ± 27.8	-143.4	103.4
Total Mean Misfit ----- (rms) = 40.8 mm				

Observed Changes are given with the conventions: Post-Seismic value -Preseismic value, and Station B - Station A.